

Elastic Properties of Polycrystalline Magnesium Carbonates to 10 Kilobars  
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 Compresstite Aggregates of magnesium carbonate  
 are reported. Velocity-density relationships similar to that observed for the olivine series. Iron substitution for magnesium in carbonates has a dependence on bulk modulus, however, changes relative velocity little with iron substitution. The  
 in carbonates affects the velocities and density and modulus decreases Poisson's ratio. The  
 bulk modulus, however, changes relatively little with iron substitution. Magnesium carbonate  
 on mean atomic weight basis decreases the shear modulus and increases Poisson's ratio. The  
 poly crystalline aggregates have similar effects on velocity and density and modulus  
 substitution of 3d transition elements for magnesium.

An important source of information on the  
 composition of the earth's interior is derived  
 from measurements of elastic-wave velocities  
 to obtain mantle compositions by estimating  
 velocities in olivine aggregates with varying iron  
 content [e.g., Birch, 1961b, 1969; Christensen,  
 1968]. Recently, new experiments with varying iron  
 content available from single-crystal studies [Gra-  
 ham and Barach, 1969; Kumazawa and Anderson,  
 1969] and from measurements of hot  
 pressed syenitic aggregates [Chung, 1970;  
 Dumitrescu et al., 1970] and naturally occurring  
 olivine dikes [Mao et al., 1971]. The results of these  
 studies agree well with compressional-wave  
 density relationships predicted by Birch [1961a]  
 and show that shear-wave velocities in olivine  
 also decrease with increasing iron content.

The samples are naturally occurring aggregate-  
 gats. Chemical analyses determined by atomic  
 absorption and expressed in weight per cent  
 are given in Table I. X-ray diffraction patterns  
 are siderite and rhodochrosite shown no addi-  
 tional minerals present in either sample. Thus  
 the moderate amounts of calcium in the rhodo-  
 chrosite and magnesite in the siderite are within  
 these and magnesium amounts of magnetite  
 the crystal latites in solid solution. A small  
 peak that was observed at a d value of 3.04 A  
 for the magnesite sample suggests that some of  
 the calcium in the chemically  
 how variations in composition within different  
 mineral groups affect elastic-wave velocities and  
 moduli. Birch [1961a] observed that, in addition to the olivine series, an isotropic line  
 connecting spline to magnetite at right angles  
 to lines of constant mean atomic weight. Re-  
 cently, Libermann [1970] has shown that the  
 density plot was approximately a vertical line  
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## SAMPLES AND DATA

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 369

TABLE 1. Chemical Analyses  
(Values in weight per cent.)

	Magnesite	Siderite	Rhodochrosite
FeCO <sub>3</sub>	0.61	86.9	0.97
MgCO <sub>3</sub>	92.6	6.68	3.05
CaCO <sub>3</sub>	6.82	1.10	7.75
MnCO <sub>3</sub>	0.18	5.14	87.2
ZnCO <sub>3</sub>		0.32	0.06
PbCO <sub>3</sub>			
Total	100.21	100.14	99.03

are given in Table 2. Barium titanate transducers of 2 MHz frequencies and ac cut quartz transducers of 1 MHz frequencies were used to generate and receive the compressional and shear waves, respectively. The data in Table 2 are uncorrected for length changes at elevated pressures, which would lower the 10-kb compressional- and shear-wave velocities by approximately 0.02 and 0.01 km/sec, respectively.

Elastic constants calculated at various pressures from the mean velocities and the densities are given in Table 3. Velocities and densities used for the calculations were corrected for dimension changes with an iterative routine and the dynamically determined compressibilities.

## DISCUSSION

Simmons [1964a, b] reported mean compressional and shear-wave velocities at 10 kb of 7.45 and 4.29 km/sec, respectively, for pure magnesite. Compressional-wave velocities were given in two directions and shear-wave velocities were given for three cores. These values are substantially lower than the mean velocities in Table 2 and are evidently due to high porosity. Simmons [1964b] reported a mean bulk density of 2.848 g/cm<sup>3</sup> for his sample, which is considerably lower than the 2.98 g/cm<sup>3</sup> density for pure magnesite [Deer et al., 1962]. No velocities have been reported for siderite or rhodochrosite.

TABLE 2. Velocities in Kilometers per Second

Rock	Density, g/cm <sup>3</sup>	Propagation Direction	Mode	Pressure, kb					
				0.4	2.0	4.0	6.0	8.0	10.0
Magnesite Chewelah, Wash.	2.972	A	$V_p$	7.90	8.15	8.27	8.34	8.39	8.43
	2.968		$V_p$	7.89	8.12	8.24	8.32	8.37	8.42
	2.966		$V_p$	7.68	8.08	8.21	8.28	8.34	8.37
	2.969*		$V_p$	7.82	8.12	8.24	8.31	8.37	8.41
	2.972	B	$V_s$	4.55	4.62	4.67	4.70	4.72	4.73
	2.968		$V_s$	4.40	4.54	4.60	4.64	4.66	4.68
	2.966		$V_s$	4.50	4.60	4.66	4.69	4.71	4.72
	2.969*		$V_s$	4.48	4.59	4.64	4.68	4.70	4.71
Siderite Roxbury, Conn.	3.752	A	$V_p$	6.74	6.97	7.08	7.13	7.17	7.20
	3.753		$V_p$	6.61	6.89	6.99	7.03	7.06	7.08
	3.756		$V_p$	6.61	6.93	7.04	7.10	7.14	7.16
	3.754*		$V_p$	6.65	6.93	7.04	7.09	7.12	7.15
	3.752	B	$V_s$	3.59	3.67	3.73	3.76	3.78	3.80
	3.753		$V_s$	3.46	3.55	3.60	3.63	3.65	3.67
	3.756		$V_s$	3.43	3.51	3.56	3.59	3.60	3.61
	3.754*		$V_s$	3.49	3.58	3.63	3.66	3.68	3.69
Rhodochrosite Catamarea Province, Argentina	3.569	A	$V_p$	6.90	7.04	7.10	7.13	7.15	7.17
	3.563		$V_p$	6.99	7.13	7.24	7.32	7.37	7.40
	3.568		$V_p$	6.70	6.85	6.93	6.98	7.01	7.04
	3.567*		$V_p$	6.86	7.01	7.09	7.14	7.17	7.20
	3.569	B	$V_s$	3.07	3.28	3.37	3.42	3.46	3.48
	3.563		$V_s$	3.46	3.49	3.53	3.56	3.59	3.61
	3.568		$V_s$	3.33	3.44	3.51	3.55	3.57	3.59
	3.567*		$V_s$	3.29	3.40	3.47	3.51	3.54	3.56

\* Mean of three preceding measurements.